

Swash Morphodynamics and Its Relation to Tsunamis

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Swash zone morphodynamics



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1. On steep beaches swash defined as the distance from point of bore collapse (max. run-down) to max. run-up.
2. Large amounts of sediment transport occur on uprush and backwash. Net transport is the difference (often small) between two large quantities.
3. Erosion / Deposition highly dynamic

Physical effects

1. Bed shear stress.
2. Turbulence / acceleration (bores)
3. Pre-suspended sediment
4. Lag effects
5. In/exfiltration
6. Pore-pressure gradients (see Tonkin et al)

A Simple Morphodynamical Model of the Swash



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Depth-integrated (NSWE) (including in/exfiltration) + Sediment Conservation & advection diffusion equation.

$$\begin{aligned}
 d_t + (dU)_x + (dV)_y &= -w \\
 (dU)_t + \left(dU^2 + \frac{gd^2}{2} \right)_x + (dUV)_y &= gdh_x - \frac{f_w}{2} U |\vec{U}| - Uw \\
 (dV)_t + (dUV)_x + \left(dV^2 + \frac{gd^2}{2} \right)_y &= gdh_y - \frac{f_w}{2} V |\vec{U}| - Vw \\
 B_t + \xi \vec{\nabla} \cdot \vec{q} + C_t &= 0 \\
 C_t + \vec{U} \cdot \vec{\nabla} C - w_s C_b - \mu_v C_{zb} &= \vec{\nabla} \cdot (\mu_h \vec{\nabla} C) - B_{st}
 \end{aligned}$$

$$\frac{1}{1-n} = \xi$$

Where free slip condition applied $B_t + \vec{U} \cdot \vec{\nabla} B - W = w$ at $z = B(\vec{x}, t)$

Timescales



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Traditional morphodynamical models decouple hydro- and morphodynamic timescales. In the swash flows are transcritical, so this is not possible. $T_m \approx T_h$

So, for tsunami simulations models must be fully coupled

A Simpler Morphodynamical Model of the Swash



$$d_t + (dU)_x + (dV)_y = -w$$

$$(dU)_t + \left(dU^2 + \frac{gd^2}{2} \right)_x + (dUV)_y = gdh_x - \frac{f_w}{2} U |\vec{U}| - Uw$$

$$(dV)_t + (dUV)_x + \left(dV^2 + \frac{gd^2}{2} \right)_y = gdh_y - \frac{f_w}{2} V |\vec{U}| - Vw$$

$$B_t + \xi \vec{\nabla} \cdot \vec{q} = 0$$

$$\frac{1}{1-n} = \xi$$

NSWE (including infiltration) + Sediment Conservation equation.

Where $\vec{q} = A |\vec{U}|^2 \vec{U} - \frac{1}{\tan \phi} |\vec{q}^*| \vec{\nabla} b$ Writing $A |\vec{U}|^2 \vec{U} = \alpha \vec{U}$

these can be combined to give

$$B_t \left(1 + \xi \frac{\alpha}{d} \right) = -\xi d \vec{U} \cdot \vec{\nabla} \frac{\alpha}{d} + \xi \frac{\alpha}{d} (\eta_t + w) + \xi \vec{\nabla} \cdot \left(\frac{\alpha}{\tan \phi} |\vec{U}| \vec{\nabla} b \right)$$

Physical Processes



Flow divergence: increasing depth
(convergence) => deposition

Infiltration => deposition

$$B_t \left(1 + \xi \frac{\alpha}{d} \right) = -\xi d \vec{U} \cdot \vec{\nabla} \left(\frac{\alpha}{d} \right) + \xi \frac{\alpha}{d} (\eta_t + w) + \xi \vec{\nabla} \cdot \left(\frac{\alpha}{\tan \phi} |\vec{U}| \vec{\nabla} b \right)$$

Flow picks up / drops based on
local concentration gradient

Diffusion (smoothing):
opposes bed growth

Cross-shore swash profile evolution



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Impermeable & permeable
beaches

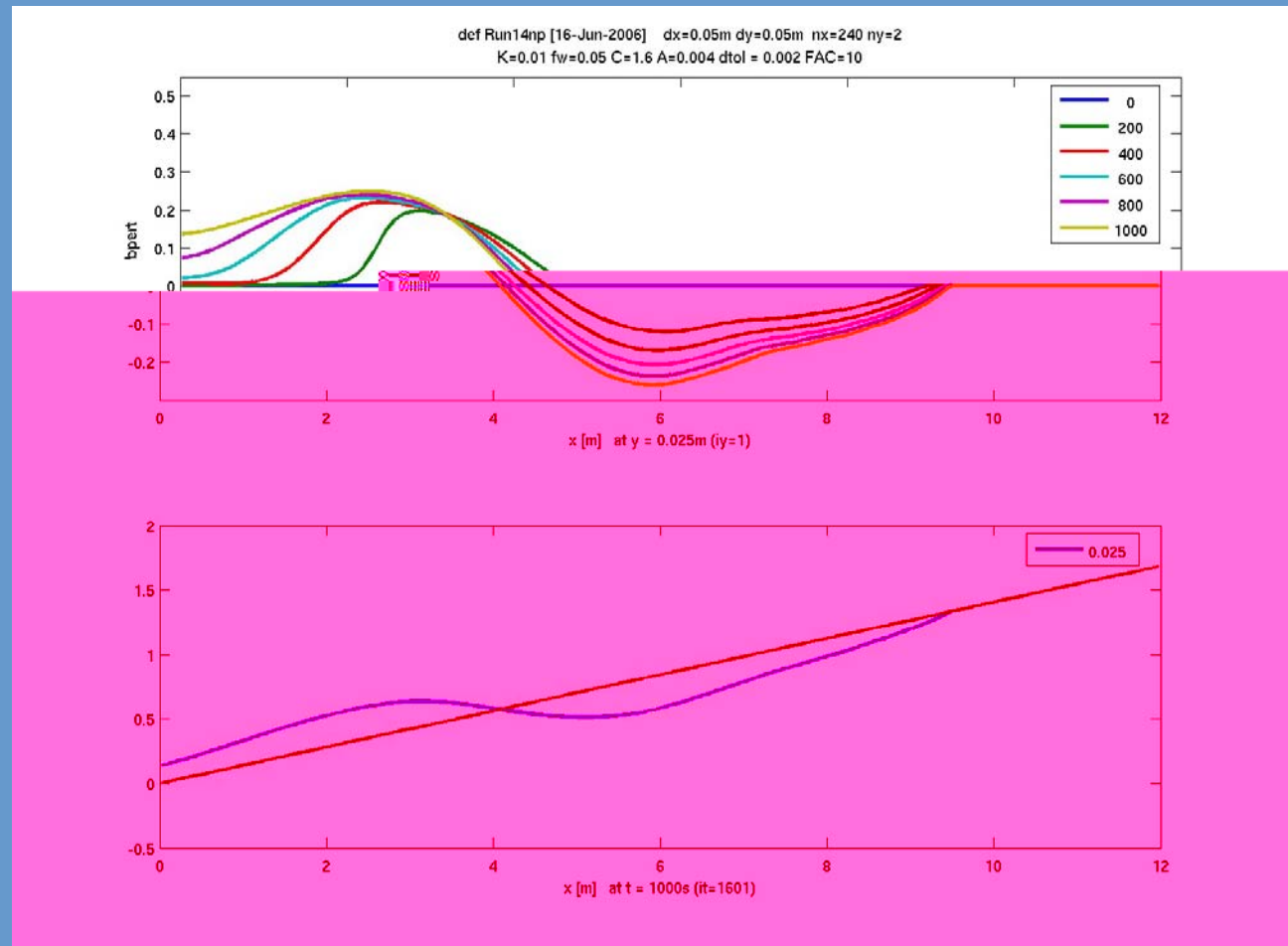
Simulations on a steep (1:7) beach,
with initially nearly non-breaking, 5s
waves

Cross-shore swash profile evolution

Impermeable beach



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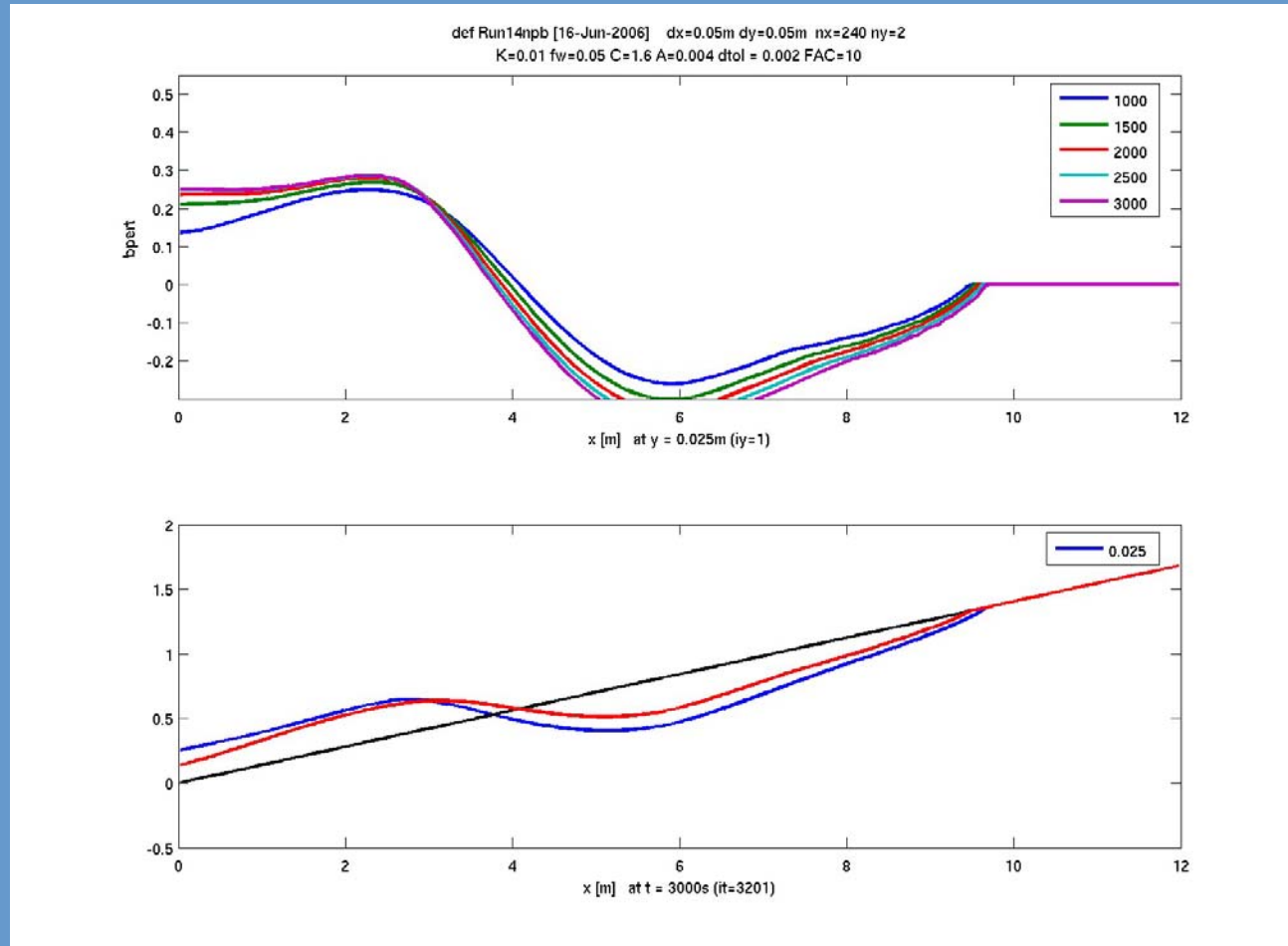
0-1000s

Cross-shore swash profile evolution

Impermeable beach



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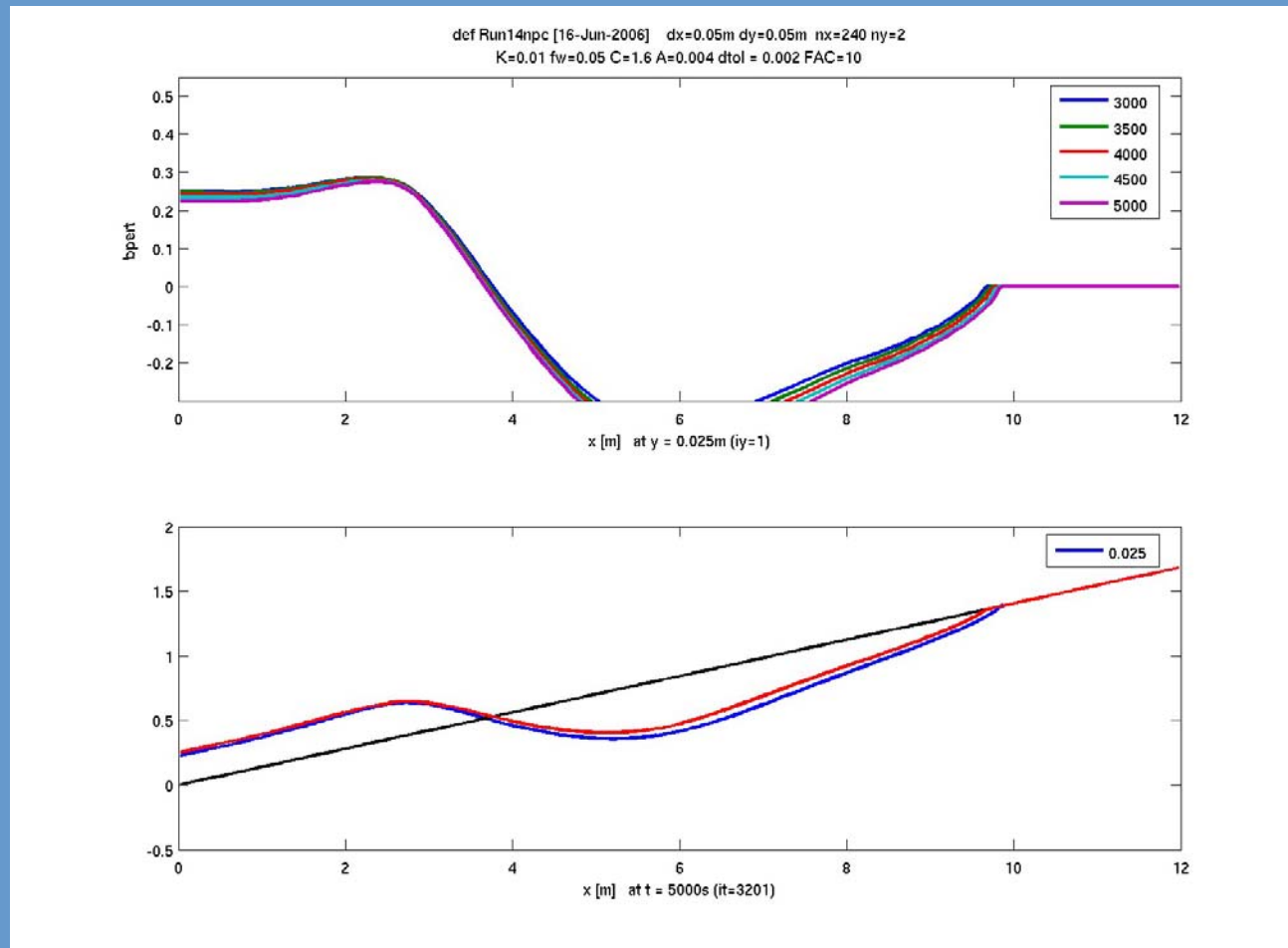
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Cross-shore swash profile evolution

Impermeable beach



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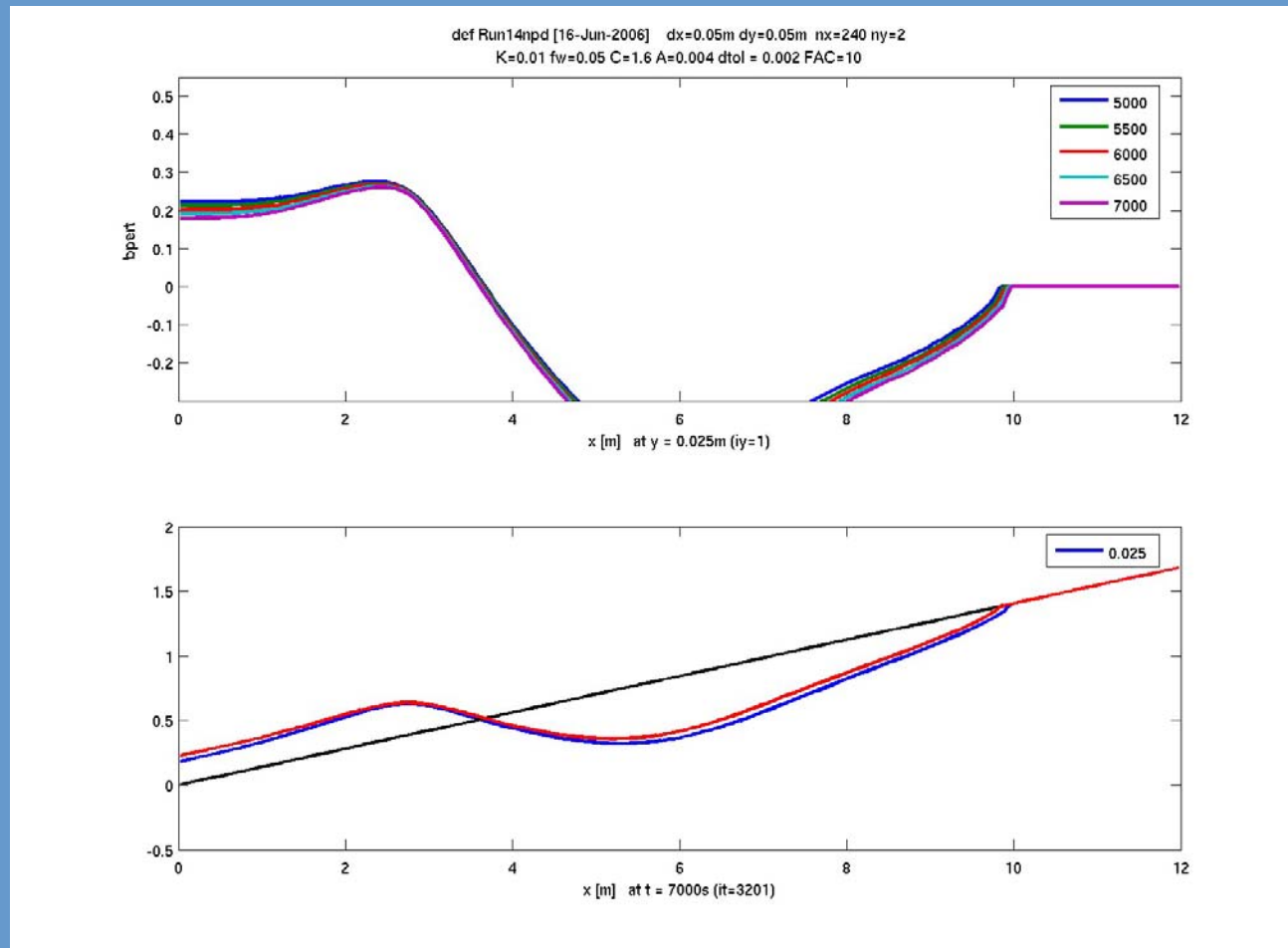
3000-5000s

Cross-shore swash profile evolution

Impermeable beach



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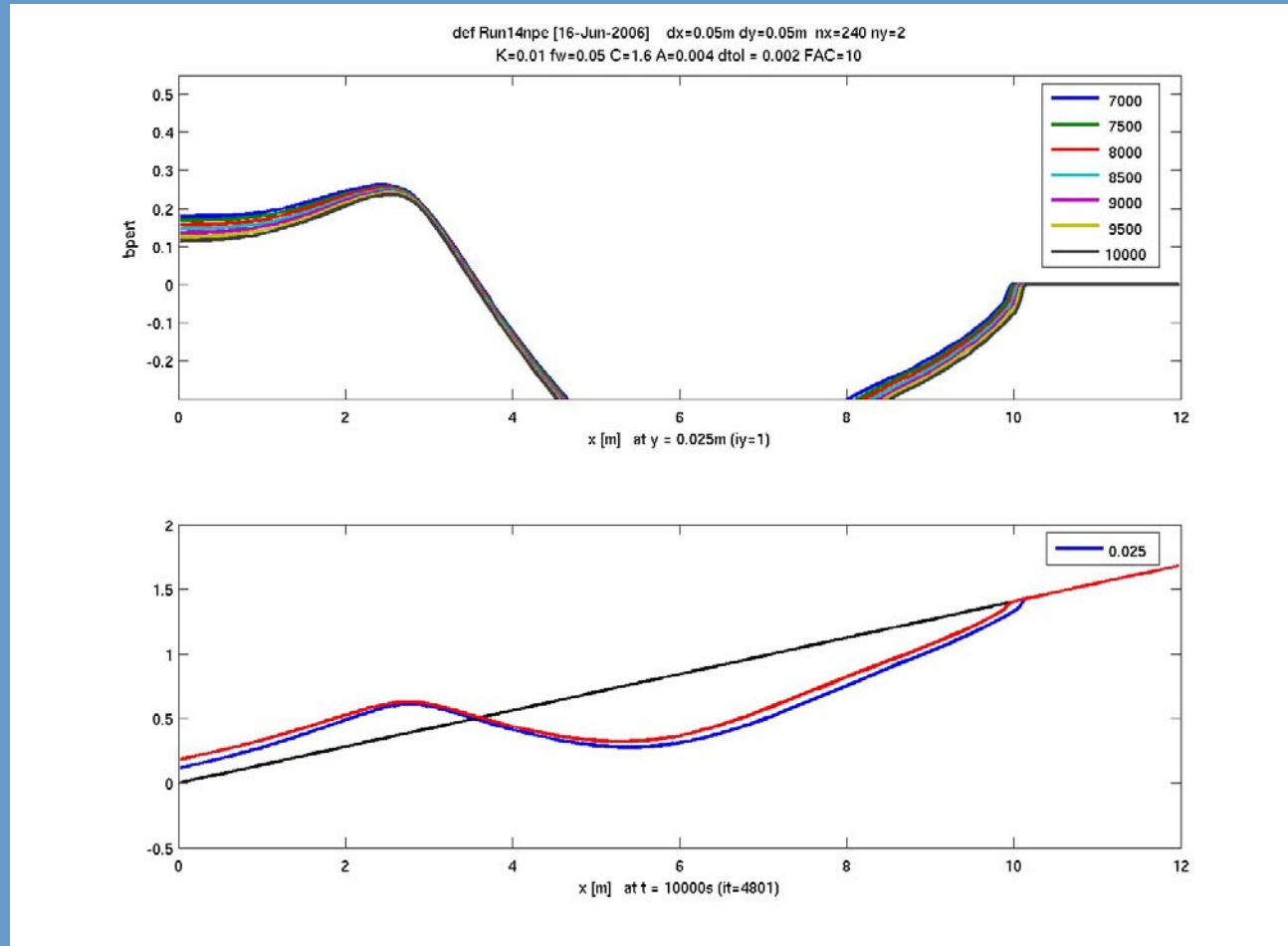
5000-7000s

Cross-shore swash profile evolution

Impermeable beach



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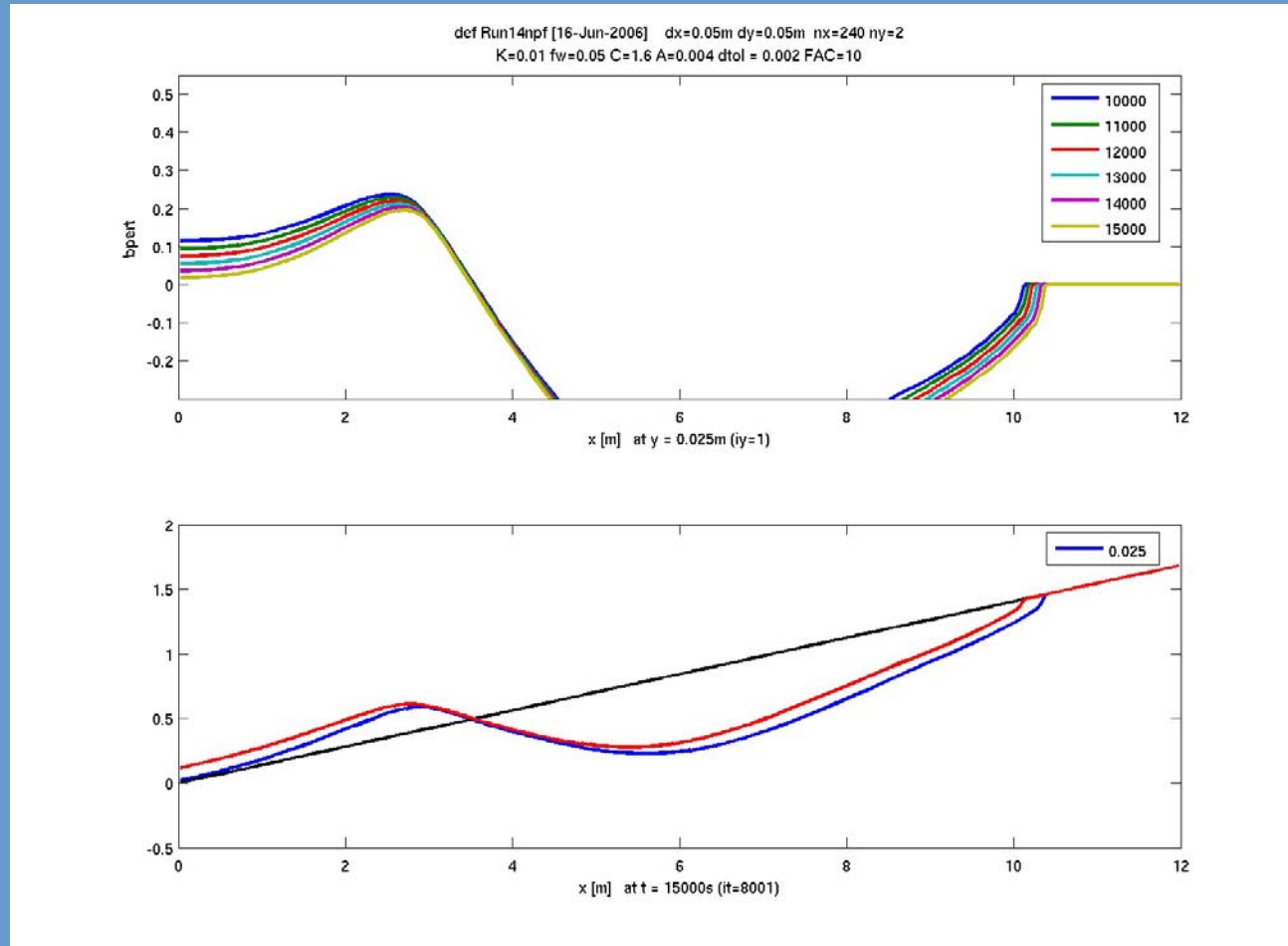
7000-10000s

Cross-shore swash profile evolution

Impermeable beach



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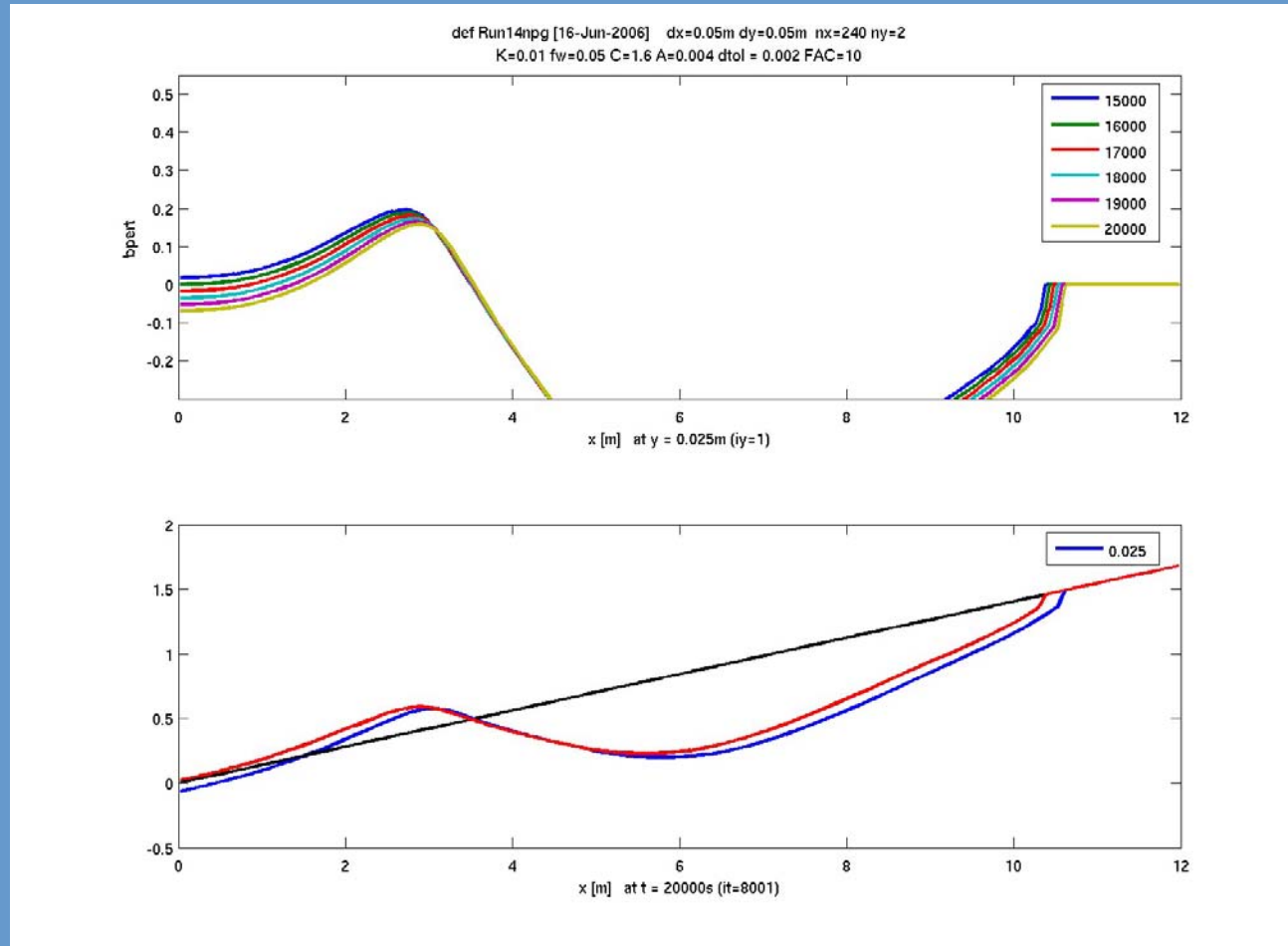
10000-15000s

Cross-shore swash profile evolution

Impermeable beach



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15000-20000s

Cross-shore swash profile evolution



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Impermeable & permeable
beaches

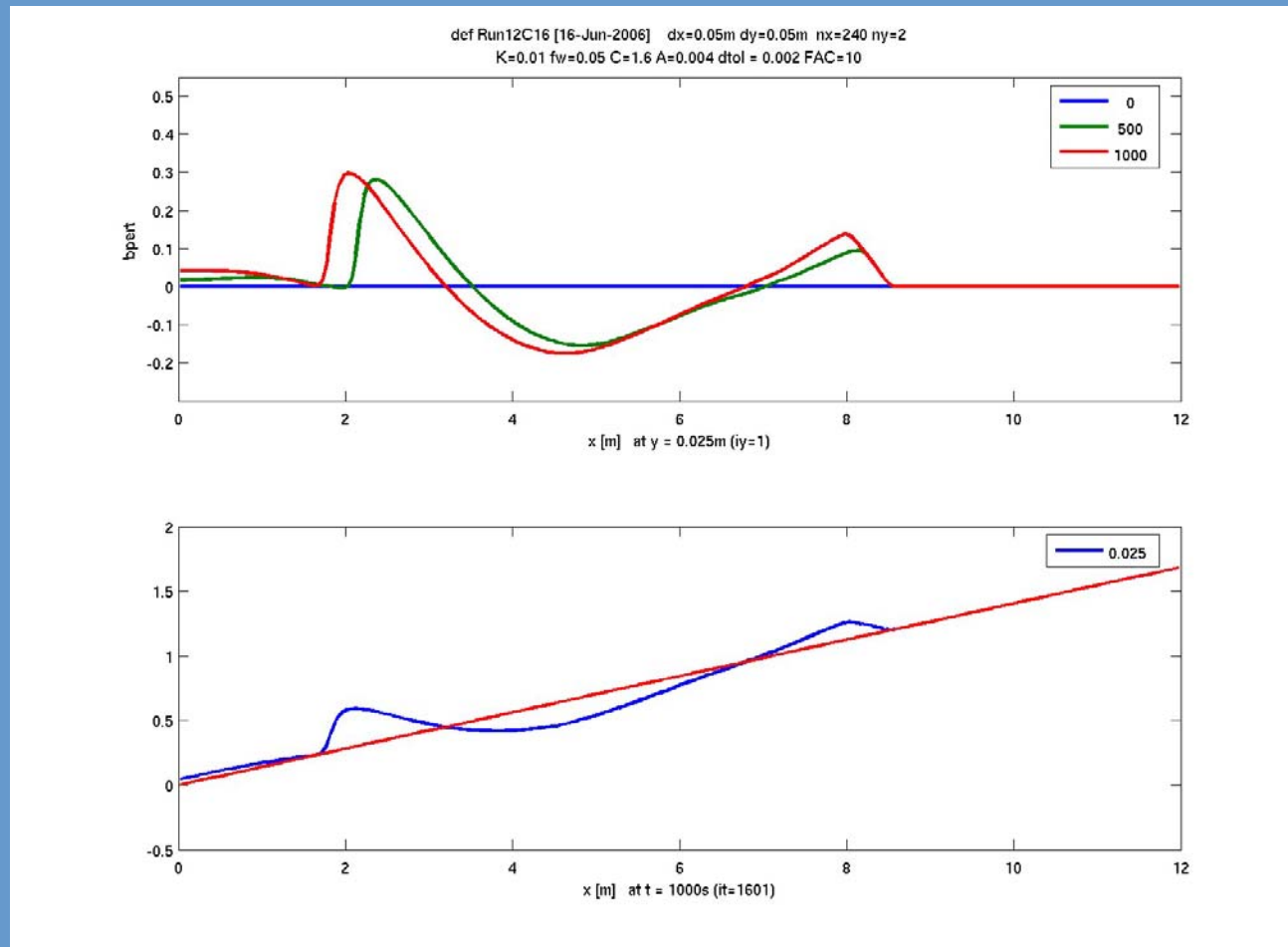
Simulations on a steep (1:7) beach,
with initially nearly non-breaking, 5s
waves

Cross-shore swash profile evolution

Permeable beach



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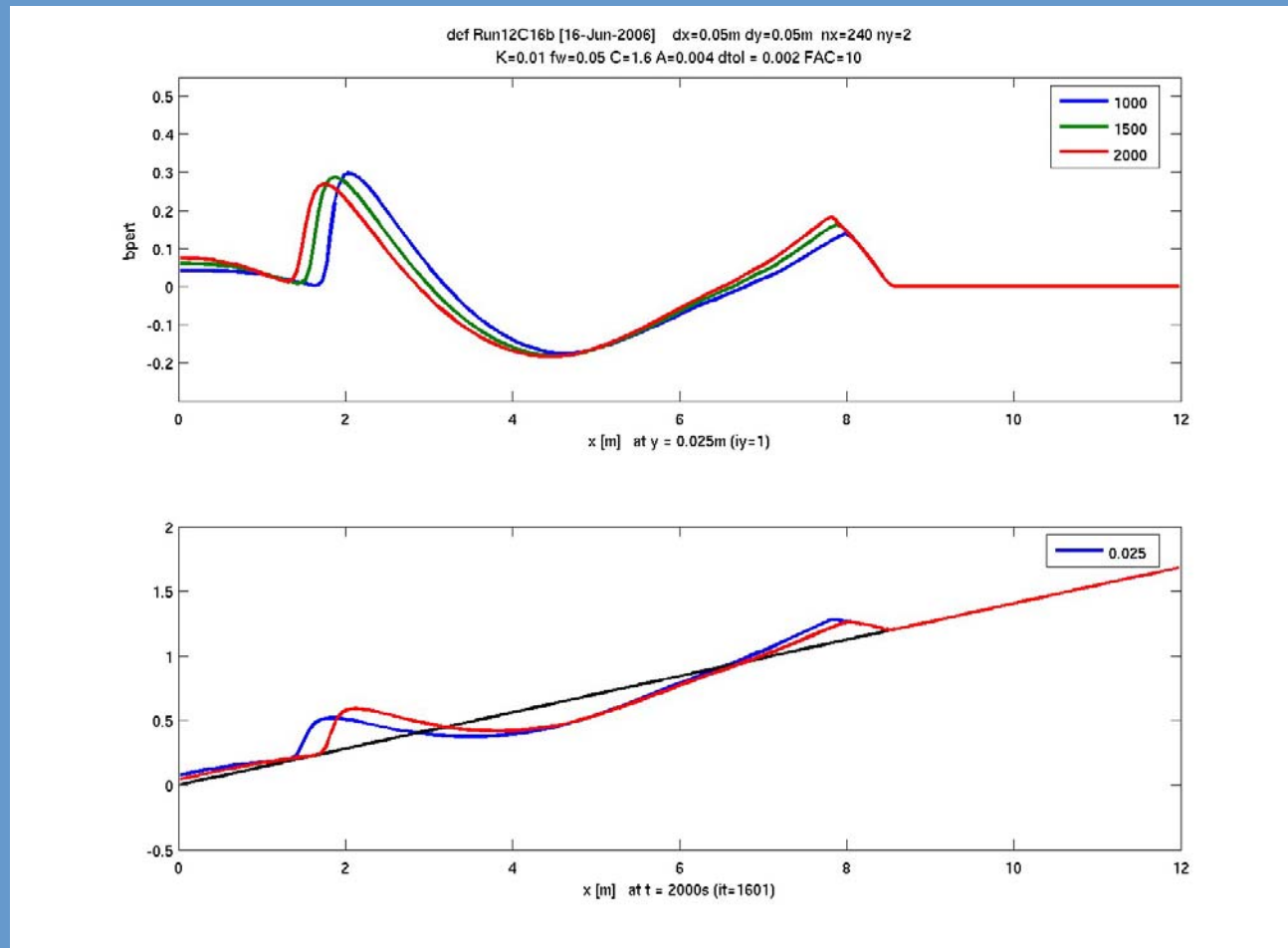
0-1000s

Cross-shore swash profile evolution

Permeable beach



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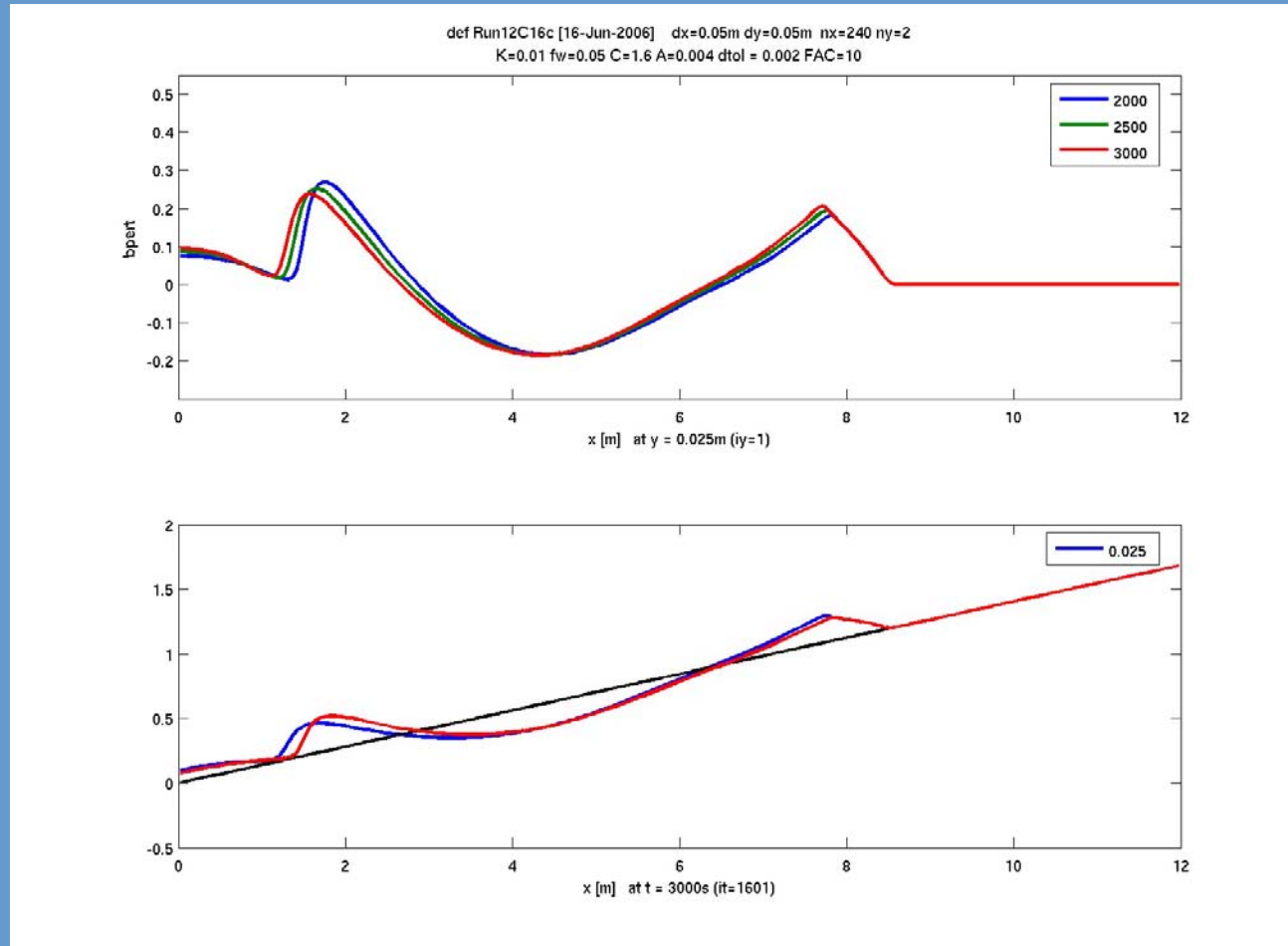
1000-2000s

Cross-shore swash profile evolution

Permeable beach



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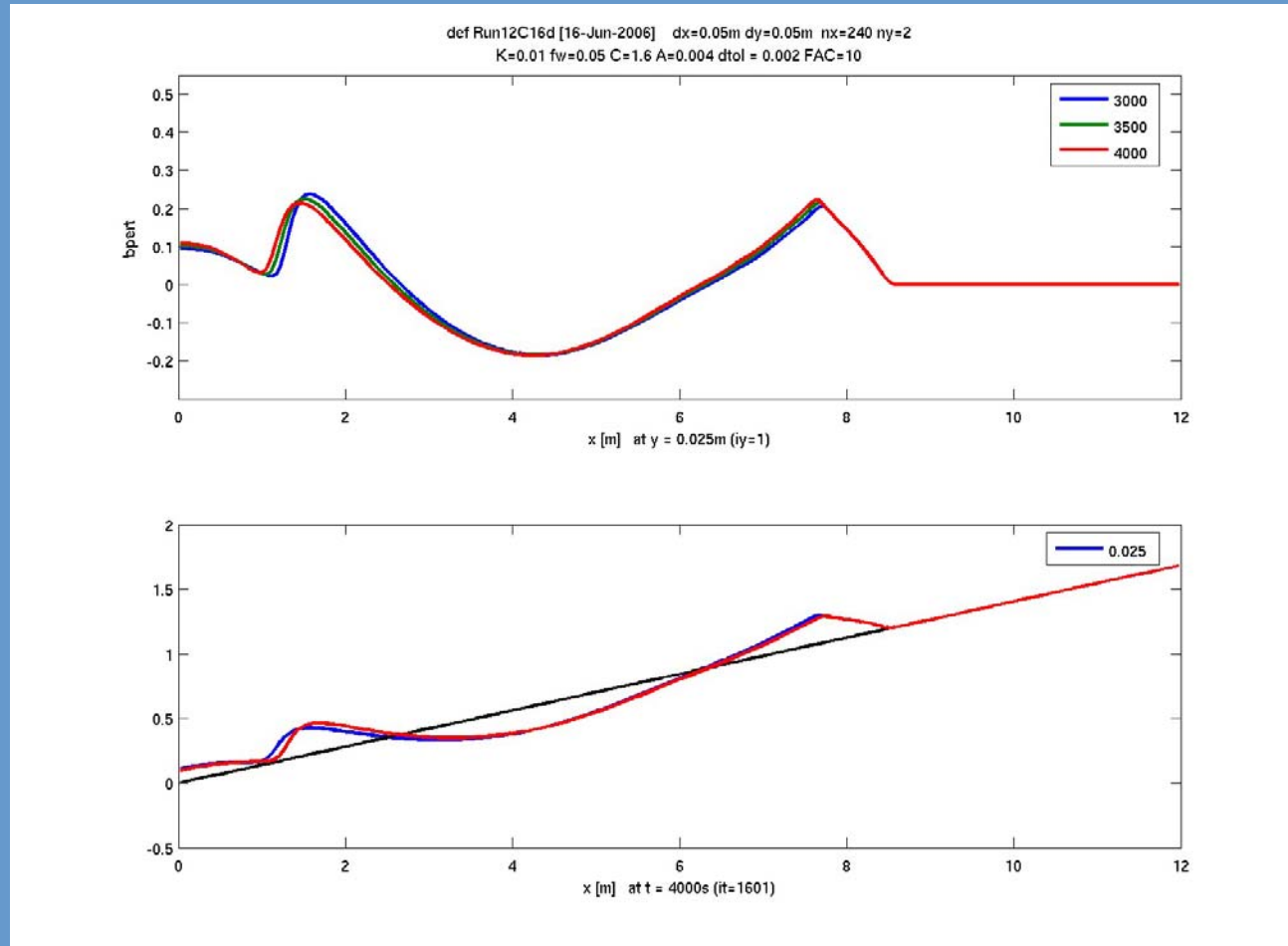
2000-3000s

Cross-shore swash profile evolution

Permeable beach



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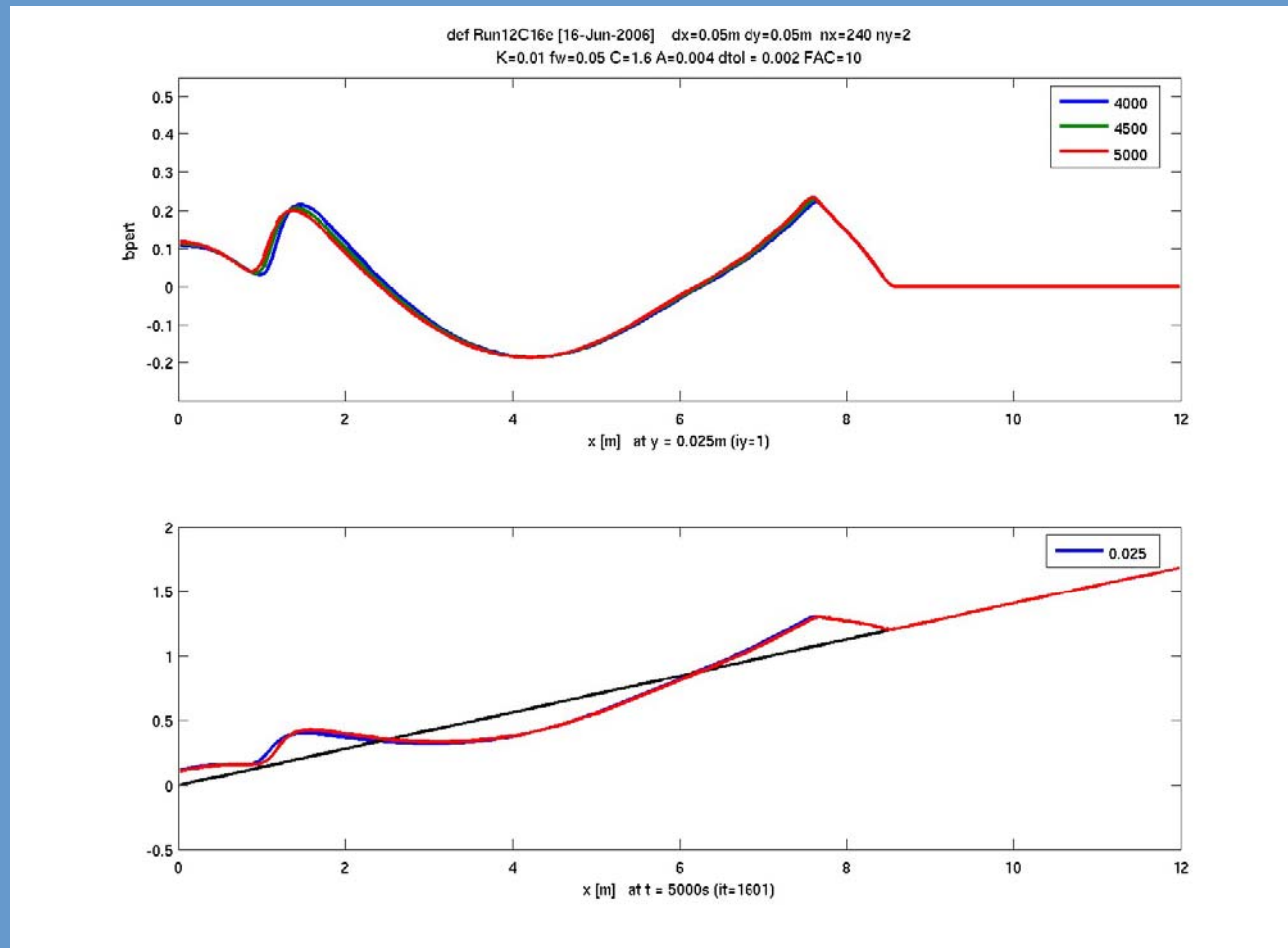
3000-4000s

Cross-shore swash profile evolution

Permeable beach



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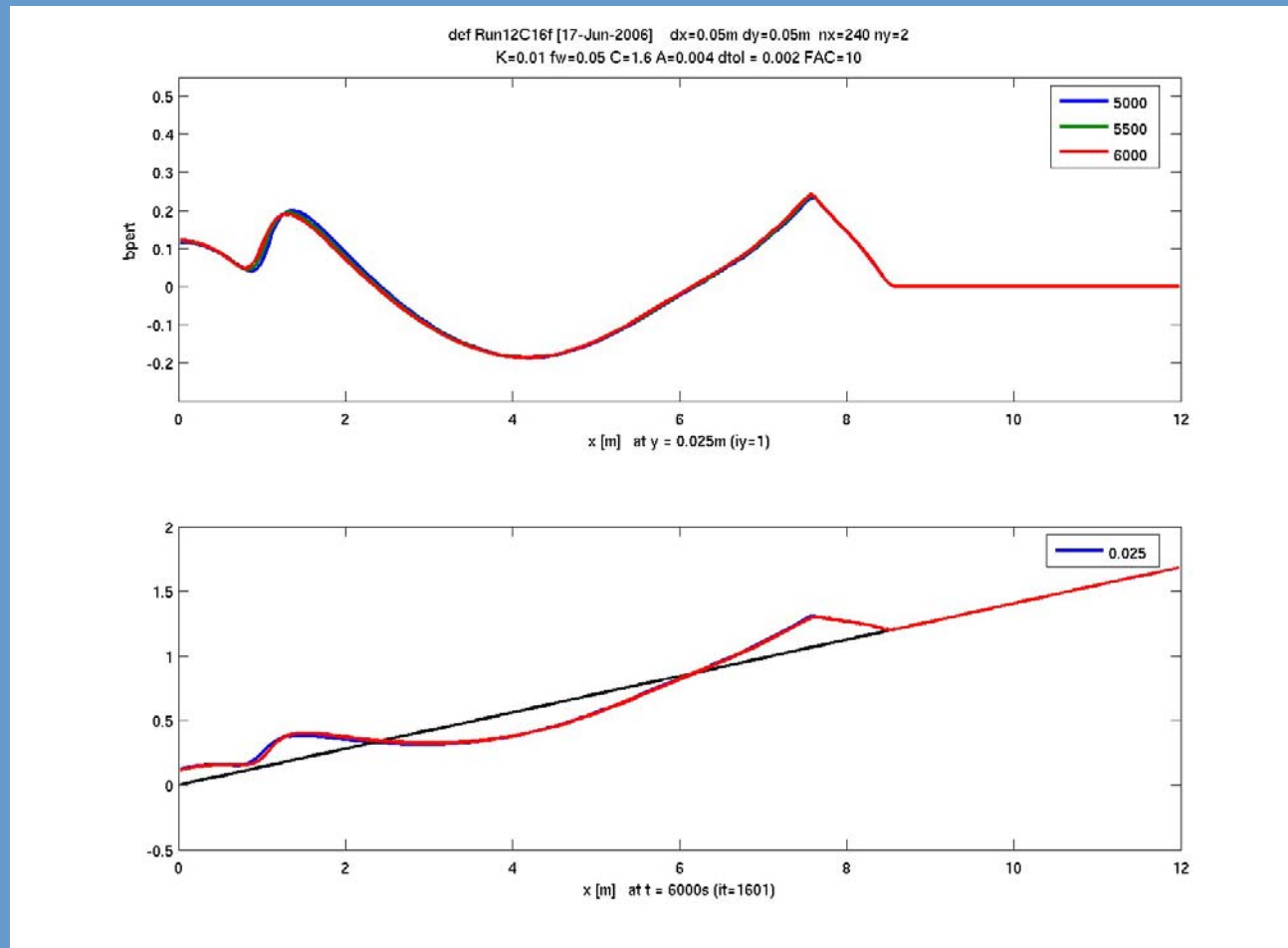
4000-5000s

Cross-shore swash profile evolution

Permeable beach



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5000-6000s

Erosion / Deposition



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1. Duration
 1. Backwash dominant in lower swash
 2. Uprush & Backwash equal in upper swash on impermeable beach—infiltration favours Uprush.
2. Uprush
 1. Erosion in lower swash (concentration gradient); deposition in upper swash (flow convergence)
3. Backwash
 1. Erosion from upper swash (flow divergence); deposition in lower swash (concentration gradient)

- These simulations are for strictly periodic waves. Phasing of tsunamis important.
- Form of tsunami at 'bore collapse' will affect subsequent depths, velocities and therefore sediment transport (e.g. Guard et al, 2006)
- Where is the upper / lower swash during a tsunami event?
Stage of tide important in determining what happens to beach

2DH Effects



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An incipient 'bump' in the upper swash will experience almost full uprush but a diminished backwash, and therefore deposition. Areas adjacent to it will experience enhanced erosion.

Infiltration will favour deposition regardless of location

- Morphological features (sand dunes) in the 'lower swash' of tsunami eroded and transported onshore.
- Lower regions in lee of morphological features might lead to weiring (accelerating flow; localised scour).
- Inundation mapping highly dependent on accurate surveying and roughness
- A lot might be learned from overtopping, fluvial flooding and dam break community